

Fluid-rock interactions to failed over-pressurisation in intrusion-hosted wallrock porphyry systems

Examples from the Northparkes district, NSW

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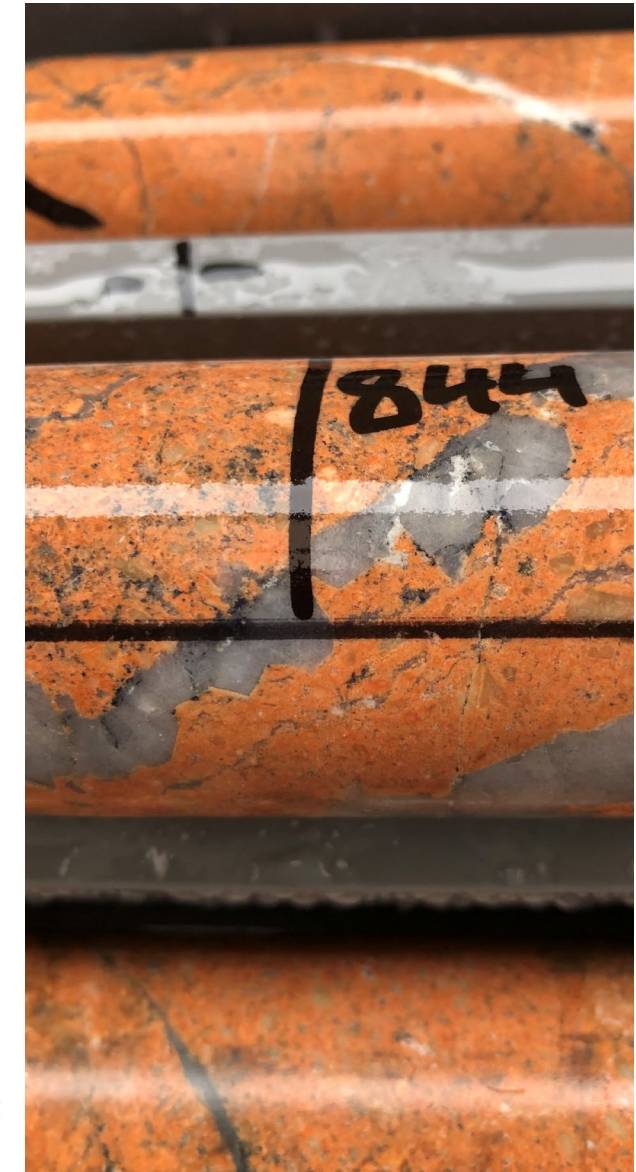
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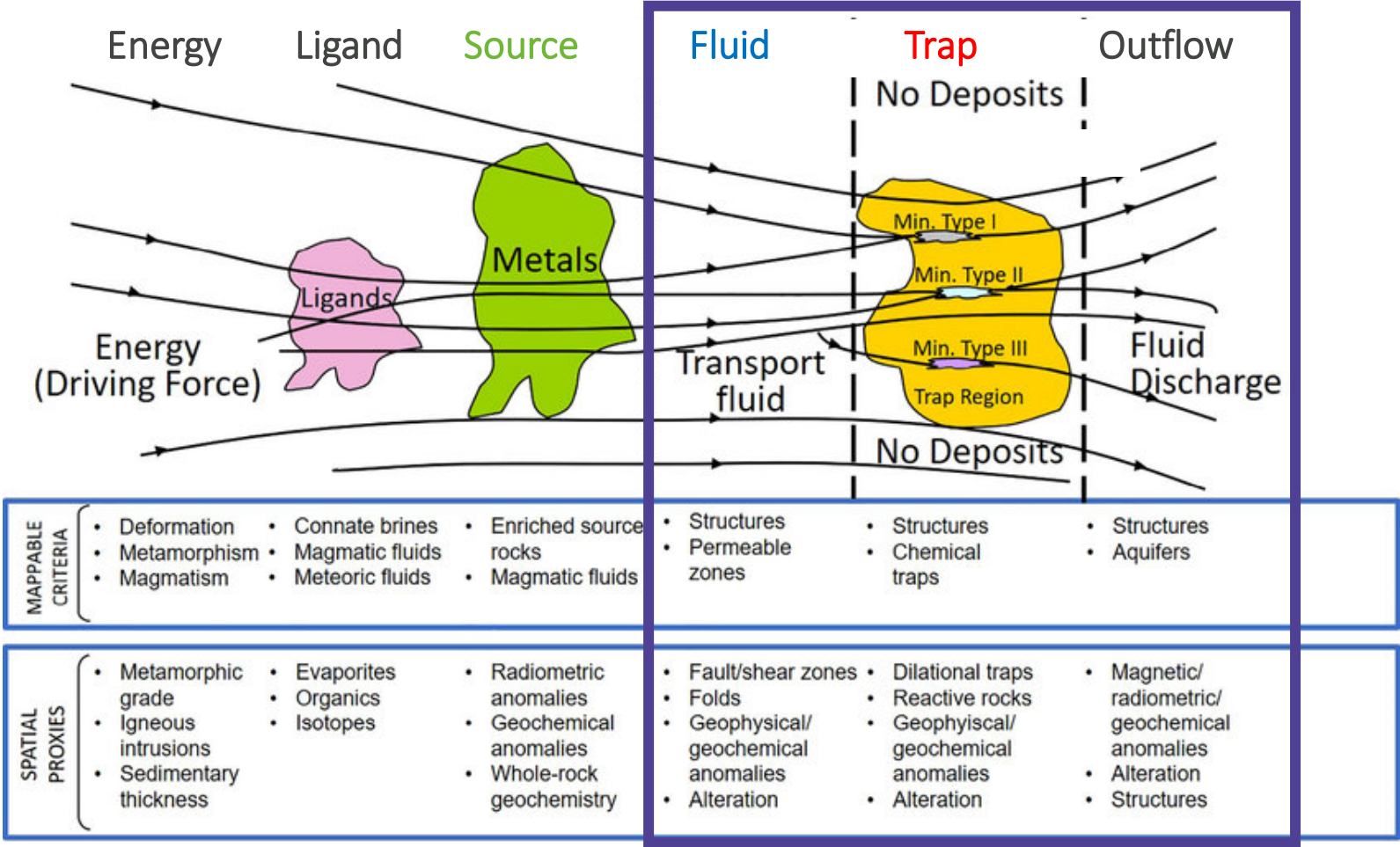
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Comb quartz crystal unidirectional solidification texture in quartz monzonite porphyry – E22 Deposit (Photo: R. Lesh)





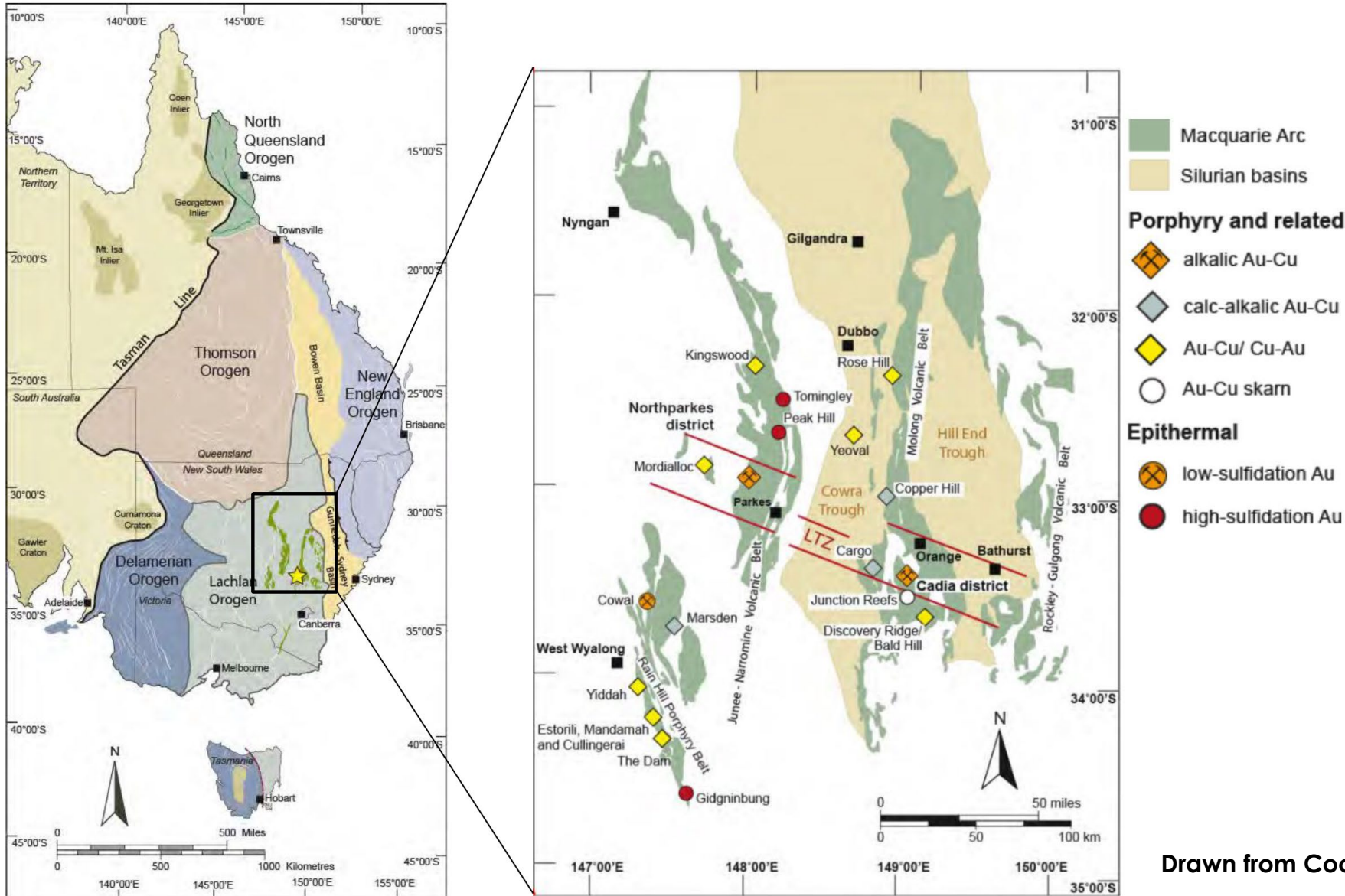
Simplified components;

- Source
- Fluid
- Trap
- Outflow

- Recent research focus on 'Source' and 'Fluid' components (magma fertility)
- 'Trap' and 'Outflow' are critical for ore formation & exploration targeting

Exploration Search Environment

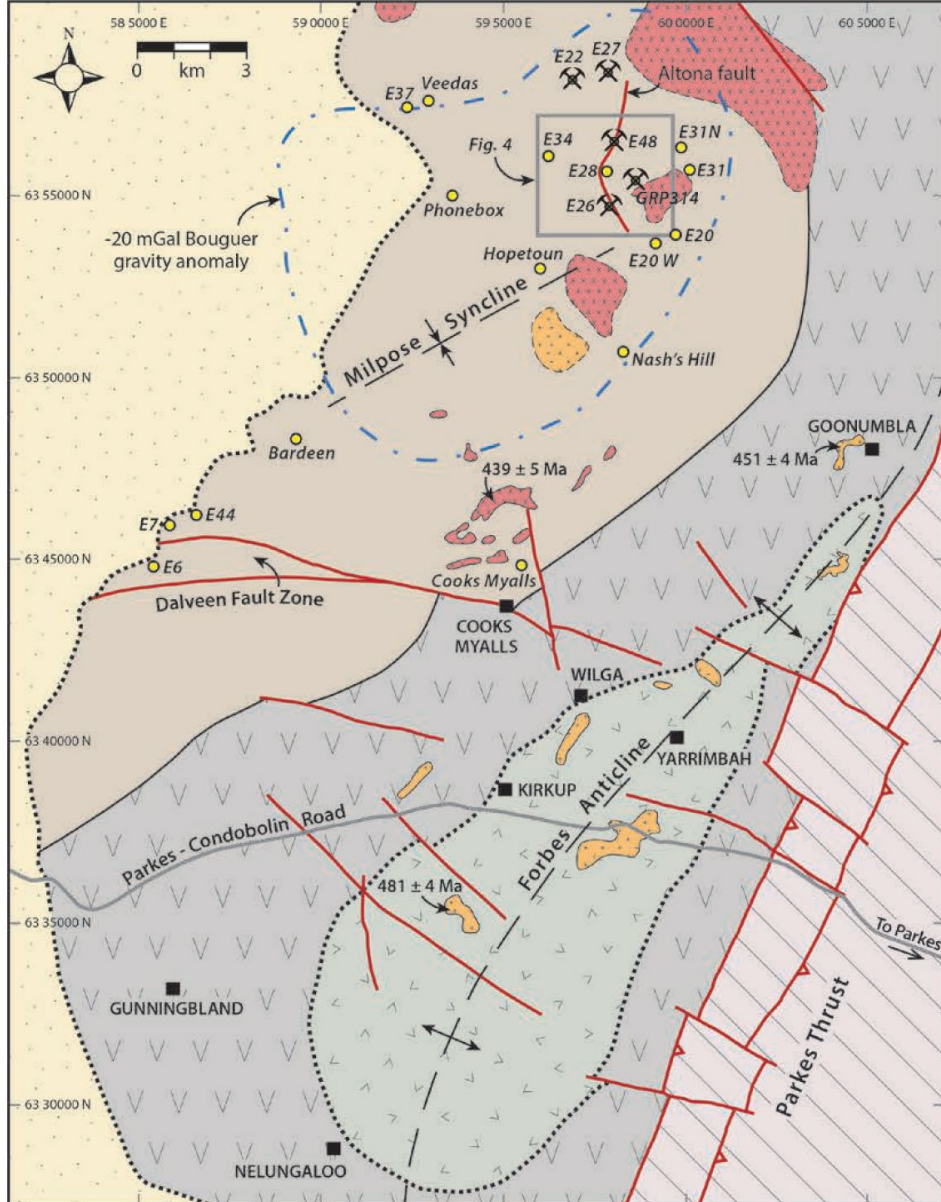
Macquarie Arc Porphyry Systems



- Segmented Ordovician – Silurian age belt
- Hosts significant alkalic porphyry Au-Cu and Cu-Au deposits
- Intrusion-hosted systems present in most districts

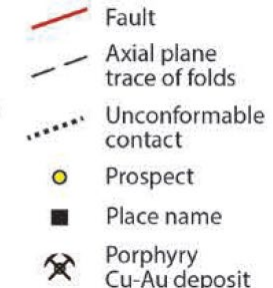
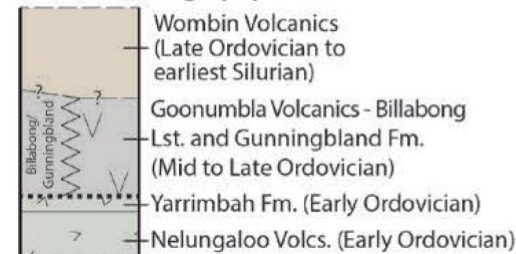
Drawn from Cooke et. al. 2009

Northparkes District Geology



- Silica-saturated alkalic Porphyry Cu-Au and related mineral systems (skarn, epithermal)
- Hosted by variable Shoshonitic to High-K Calc-Alkaline volcanic and volcaniclastic package and comagmatic intrusive suite
- Mineralised porphyry deposits in Northparkes camp L. Ordovician to E. Silurian – 437-439Ma (Lickfold *et. al.* 2007)

Volcanic stratigraphy



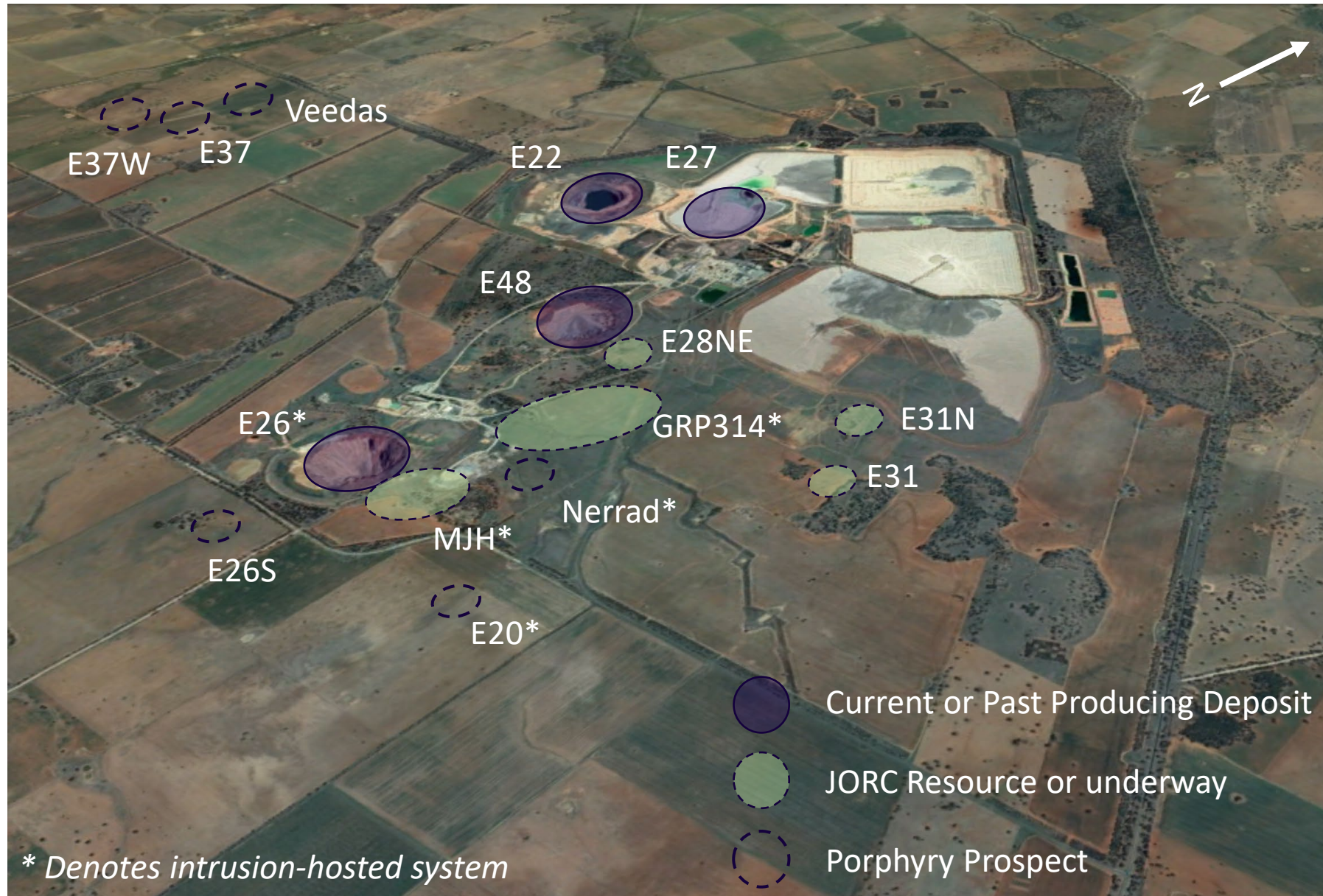
District map modified from Owens *et. al.* (2017) after Pacey (2016), Simpson *et al.* (2005), Arundell (1998) and Heithersay *et al.* (1990)

- 22 Porphyry systems defined to date
- Four main ore systems mined, five planned
- Discrete ore system footprints: <500m laterally
- Vertically extensive: > 1200m deep
- Mineralisation associated with potassic, sodic, and phyllic alteration around monzonite intrusions
- Zoned sulphides: bn-cpy+/-cct-tn-cov, distal py
- Generally have upright pipe-like geometries
- **Also larger, lower-grade intrusion hosted systems**
- Vein hosted, vein related, and disseminated mineralisation

E27 Open Pit – quartz monzonite intruding trachyandesite lavas and fragmental volcanics



Metal Endowment



Past Production

146.3 Mt @ 0.83% Cu,
0.30g/t Au

1.2 Mt Cu, 1.4 Moz Au

Current Reserves & Resources

607.4Mt @ 0.55% Cu,
0.21g/t Au

3.4Mt Cu, 4.0Moz Au

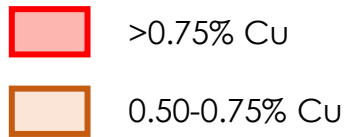
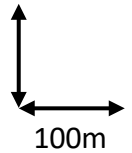
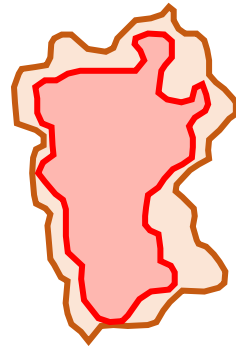
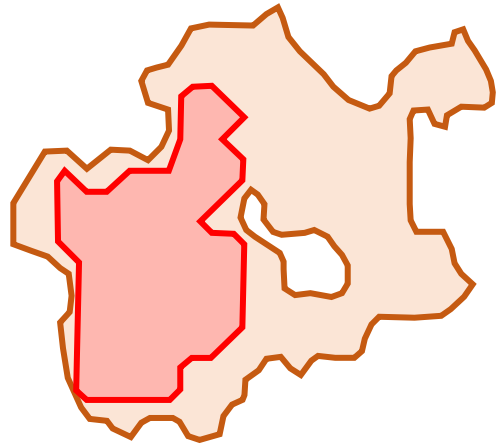
Total Mineral Endowment

~4.5Mt Cu, 5.4 Moz Au

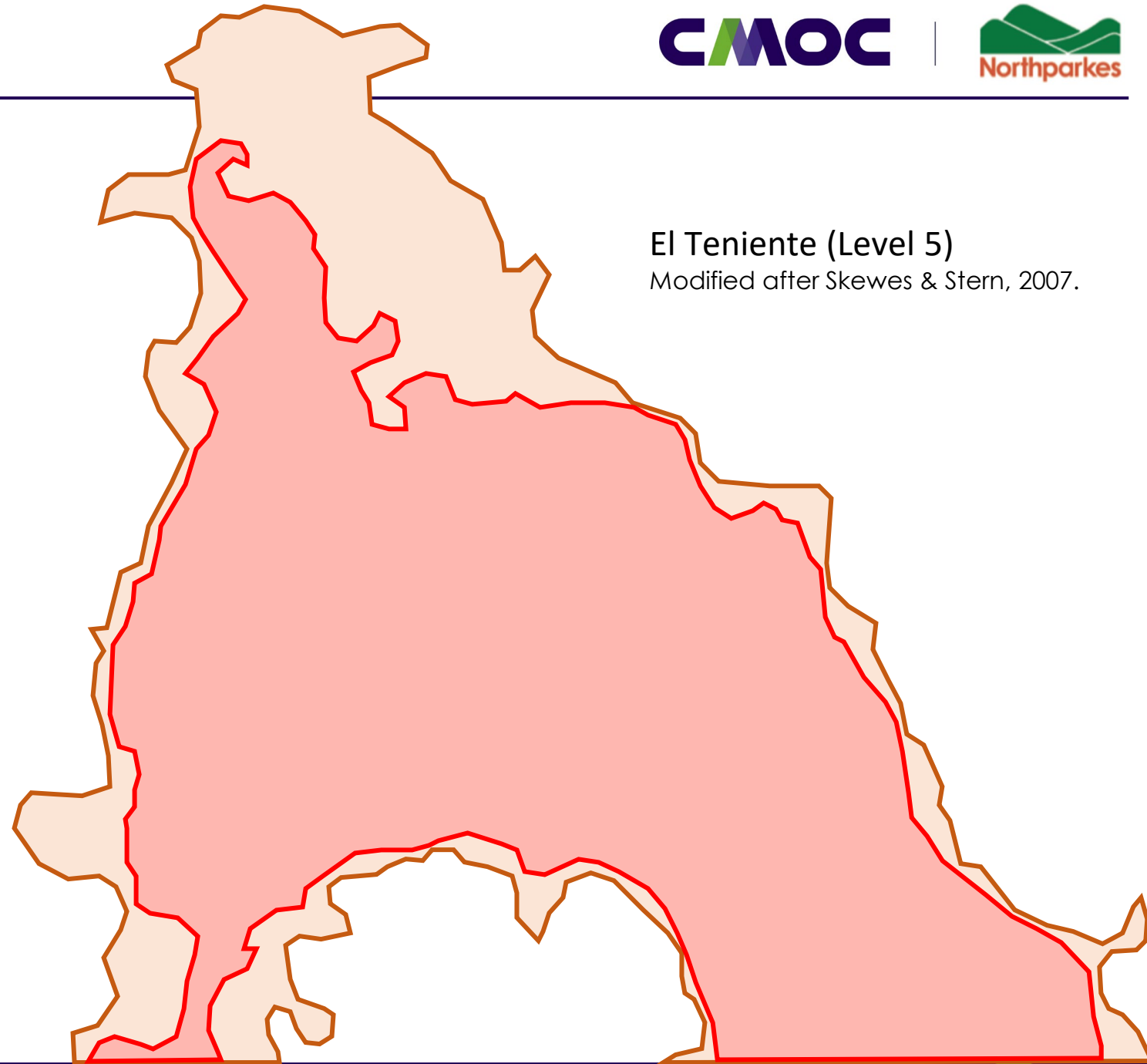
Deposit Footprints

E26
(9480RL)

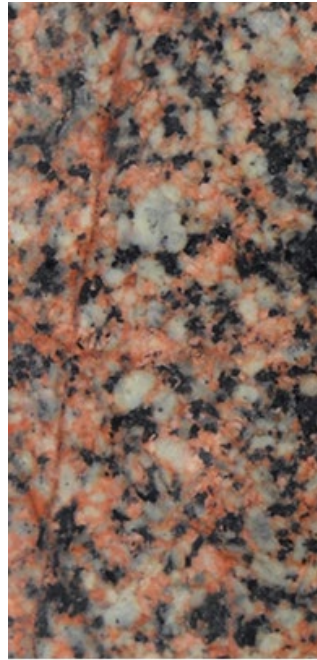
E48
(9800RL)



El Teniente (Level 5)
Modified after Skewes & Stern, 2007.



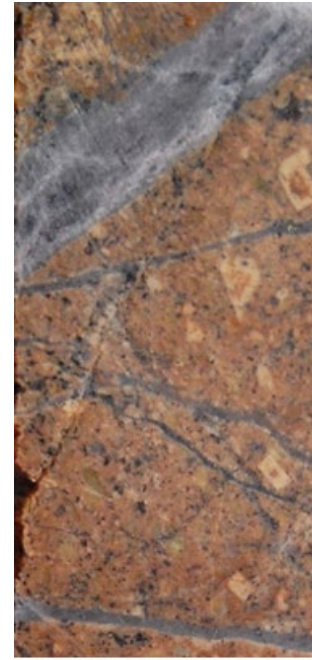
- Protracted intrusive history – 445-436Ma from Re-Os, U-Pb dating (Lickfold et al. 2007, Wells et al. 2020)
- Pre- to early-syn mineral Biotite Quartz Monzonite
- Early syn-mineral Quartz Monzonite (QMZ)
- Syn-Mineral Quartz Monzonite Porphyry (QMP)
- Late Syn-Mineral QMP, Post Mineral Monzonite Porphyry (Zero) & Late Basic Dykes (BAD)



Pre-Mineral Monzonite



Syn-min Qtz Monzonite



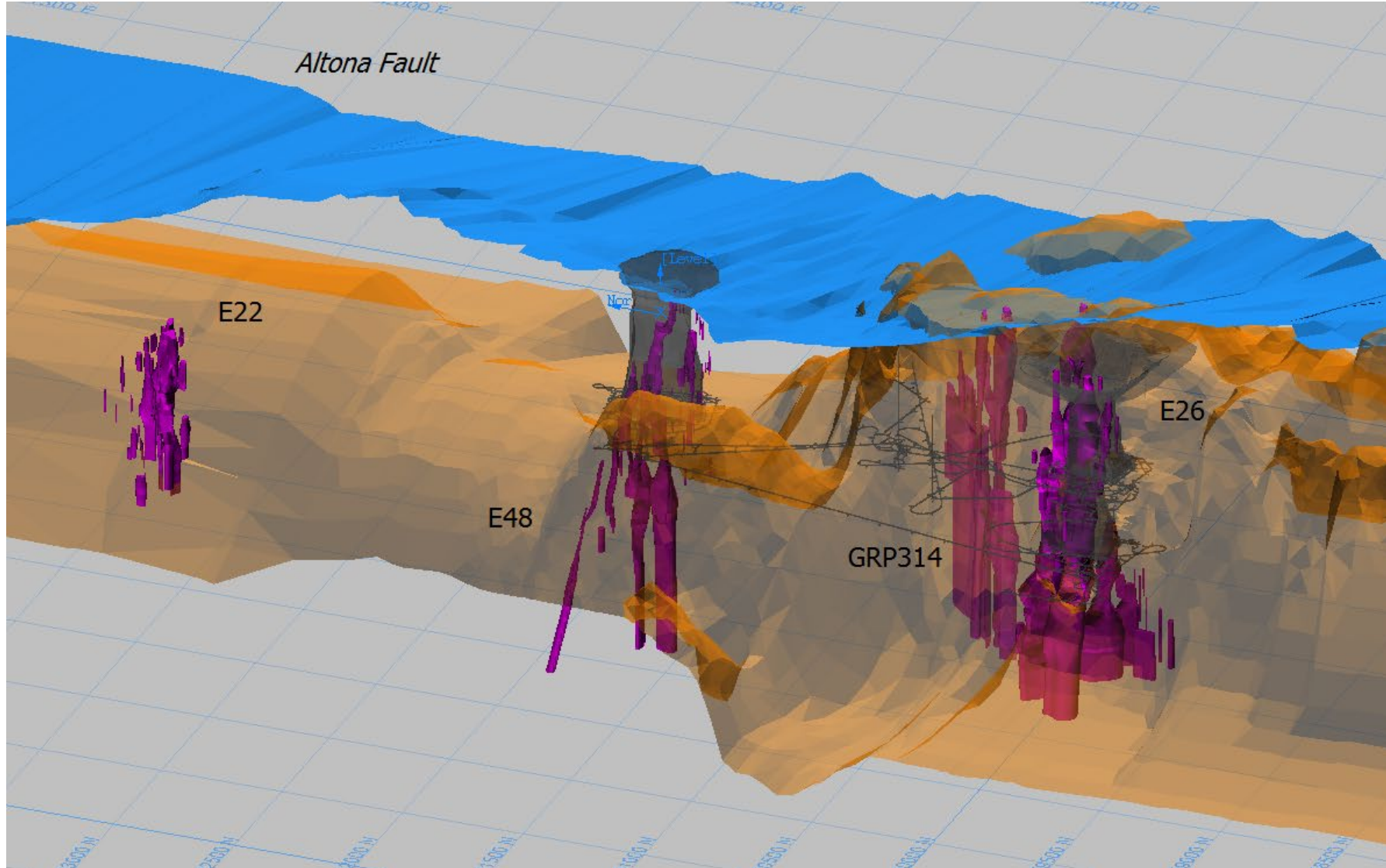
Mineralised Porphyries



Zero Porphyry

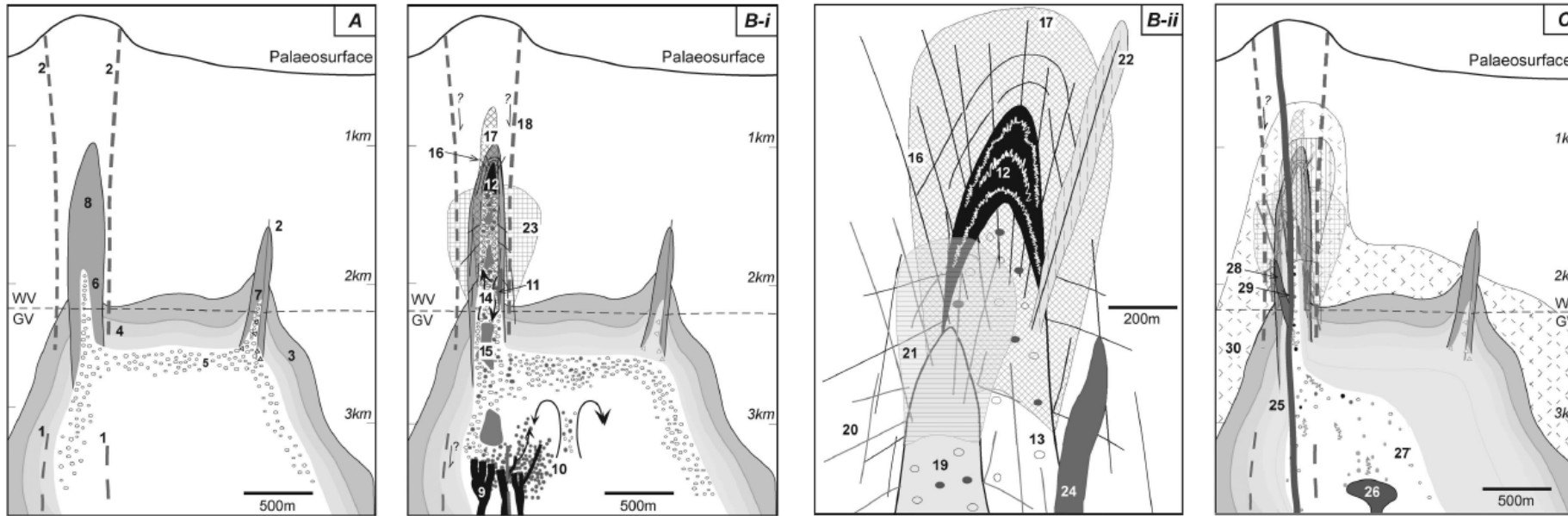


Late Basic Dyke



- All known deposits within 500m of BQM contact – fundamental control
- Syn-mineral quartz monzonite bodies intrude through BQM along subvertical structures formed during deflation of BQM (cooling and fluid loss)

Genetic Model for Porphyry Formation



Drawn from Lickfold, 2007

- A Formation of BQM magma chamber & generation of planes of weakness by shrinkage during cooling of BQM
- B-i QMP formation at depth, passing through BQM and development of mineralisation
- B-ii Successive QMP phases of decreasing volume, overprinting earlier phases and associated alteration and mineralisation, intrusion of post-mineral phases
- C Late-stage telescoping of distal alteration assemblages onto system – fault controlled

- No appreciable difference in timing of main stage mineralising monzonite intrusions (Wells et. al. 2020)
- Differences in oxidation and volume of fertile melt
- Spatial extent of sulphide mineralisation shells and ore zones in each system observed to be a function of;
 - Primary texture and grainsize of host-rock sequence (porosity and permeability)
 - Lesser (local) influence of secondary permeability caused by fluid overpressure-induced fracturing
- Permeability effects enhanced in intrusion-hosted systems - generally tight, lower grade core, with broad gradation to background copper values



Bornite rimming chalcopyrite within compositionally zoned K-feldspar megacryst in quartz monzonite porphyry – E22 Deposit (Photo: R. Lesh)

A man in a brown cowboy hat and a green shirt is holding a long wooden staff. In the background, another man in a striped shirt is visible. The scene is outdoors with green trees and a blue sky.

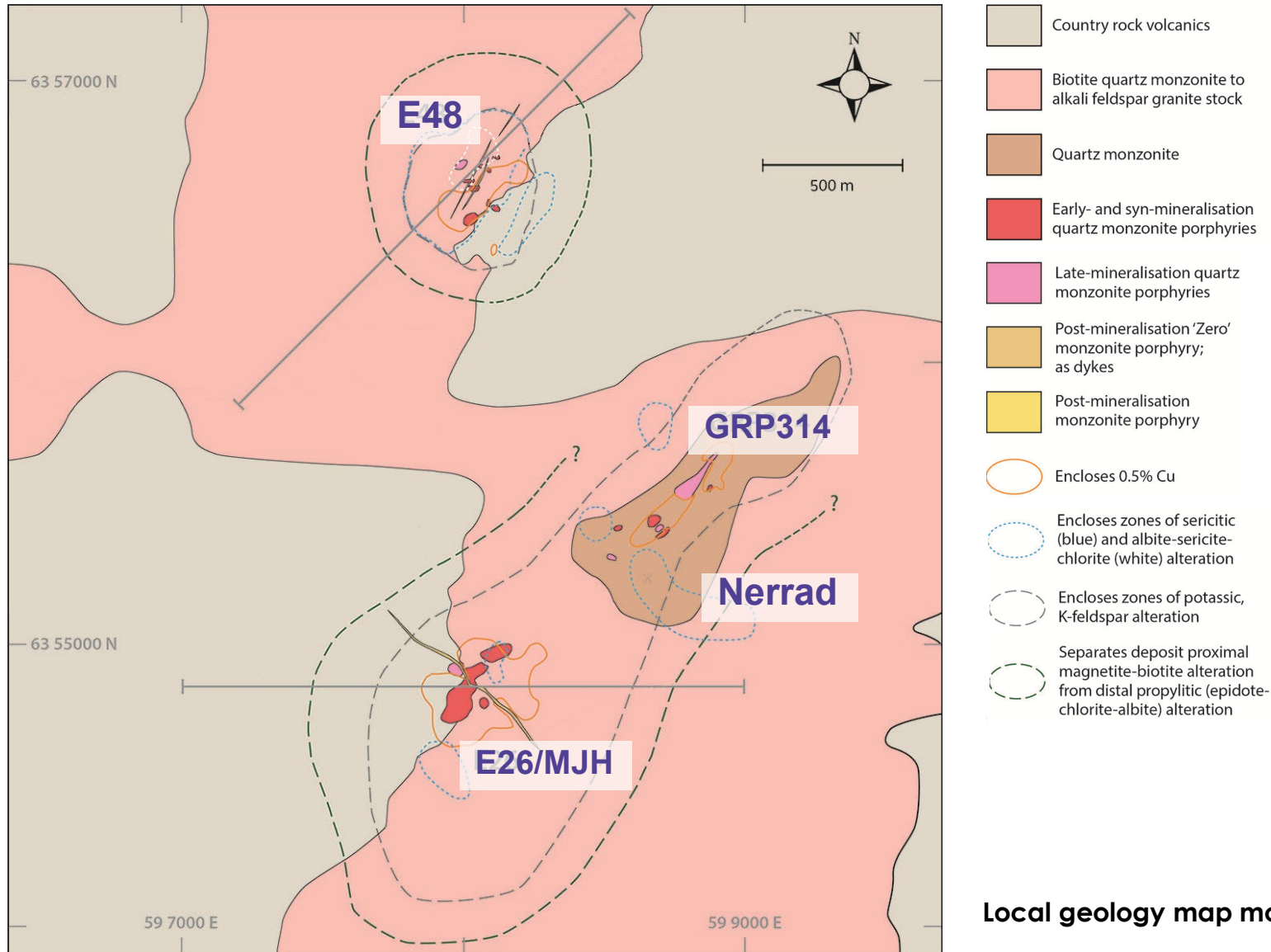
THE GOOD AND THE BAD AND THE UGLY

Not As Good

Still OK

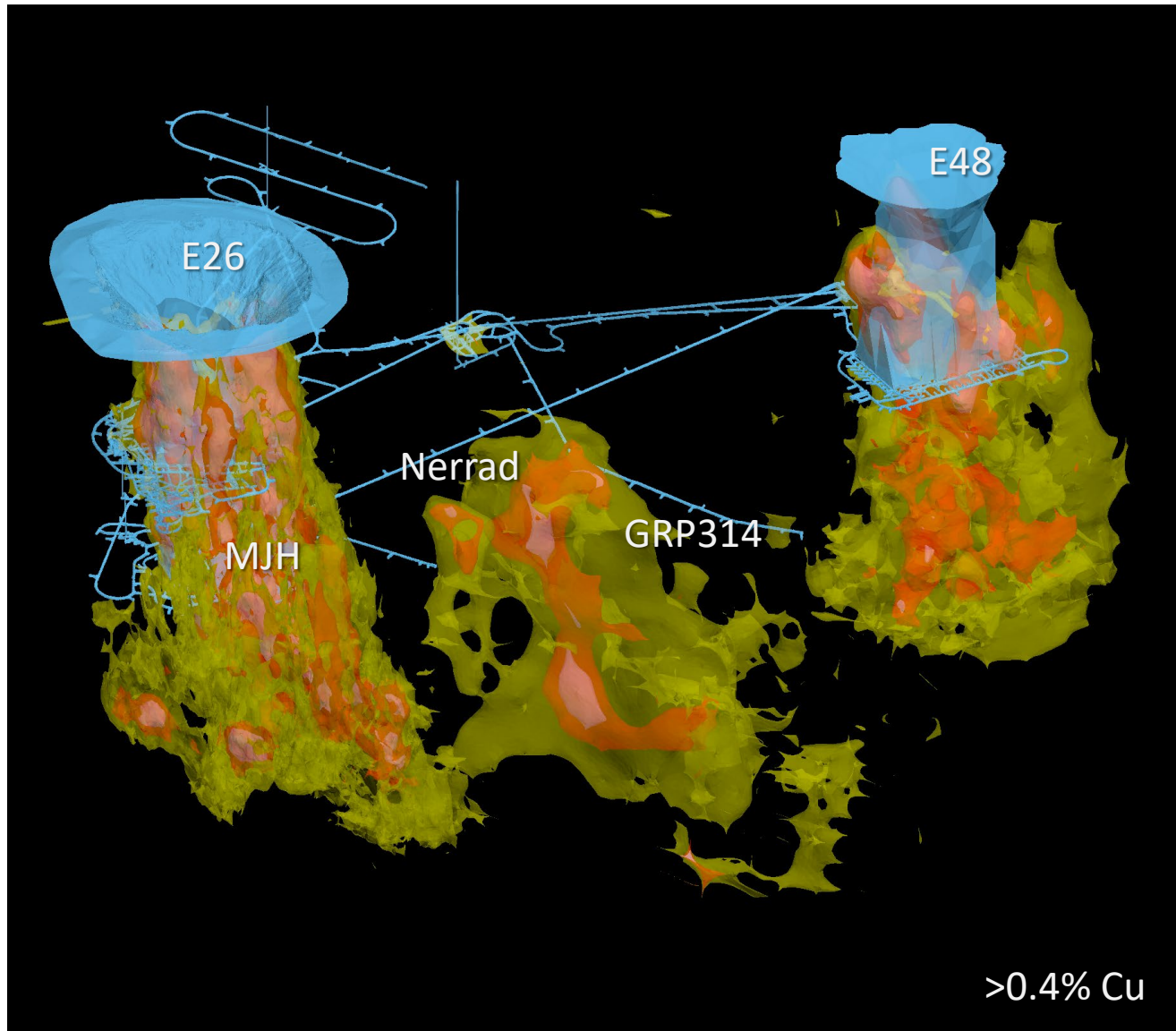
Lets talk overpressure and wall rock influence on porphyry formation in intrusion-hosted deposits

Intrusion-hosted Porphyries



- GRP314, MJH, and Nerrad entirely within pre- to early syn-mineral monzonites
- E26 and E48 mineralisation within same monzonites at depth
- GRP314 and Nerrad truncated by E-dipping low angle structure (Altona Fault)

Local geology map modified from Owens et. al. (2017) after Pacey (2016)



- Commonly broader low-grade halo compared to volcanic-hosted systems
- Ubiquitous 'red-rock' alteration
albite > k-feldspar > hematite
- Common upward-flaring structurally controlled zones of phyllic alteration
- Low As (<10ppm), elevated Zn (60-100ppm) compared to volcanic hosted systems
- Better grades (Cu) where overpressure has occurred on mineralising porphyry carapace
- Poorly defined contacts to intrusive phases

Fundamental Controls

- Fertility of the Magma (can it carry metal?)
- Volume of fertile magma you can flux up (how much metal-rich magma?)
- An efficient (but not too efficient!) plumbing system to get the magma into the shallow crust

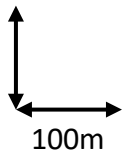
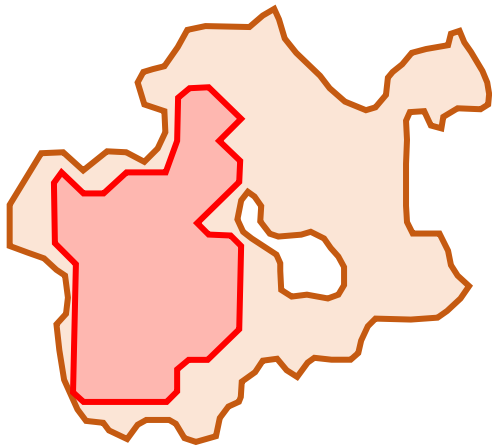
Local Factors

- The environment the magma is fluxed into in the shallow crust (porosity/permeability)
- The rate at which the magma cools and exsolves the metals (and the things that carry them...)

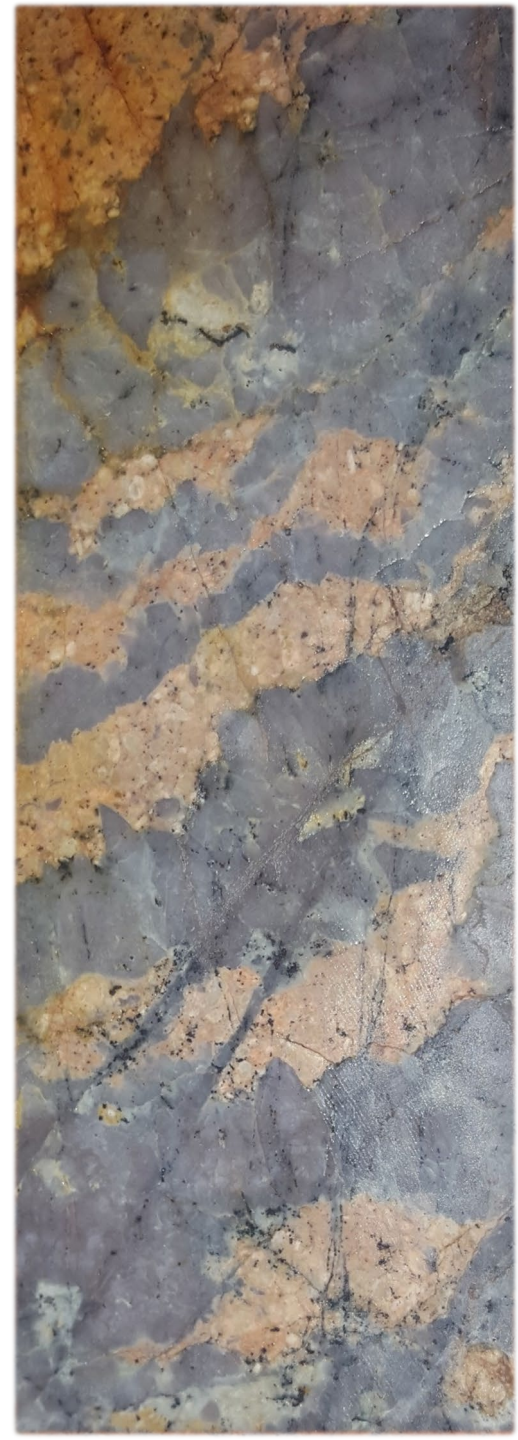


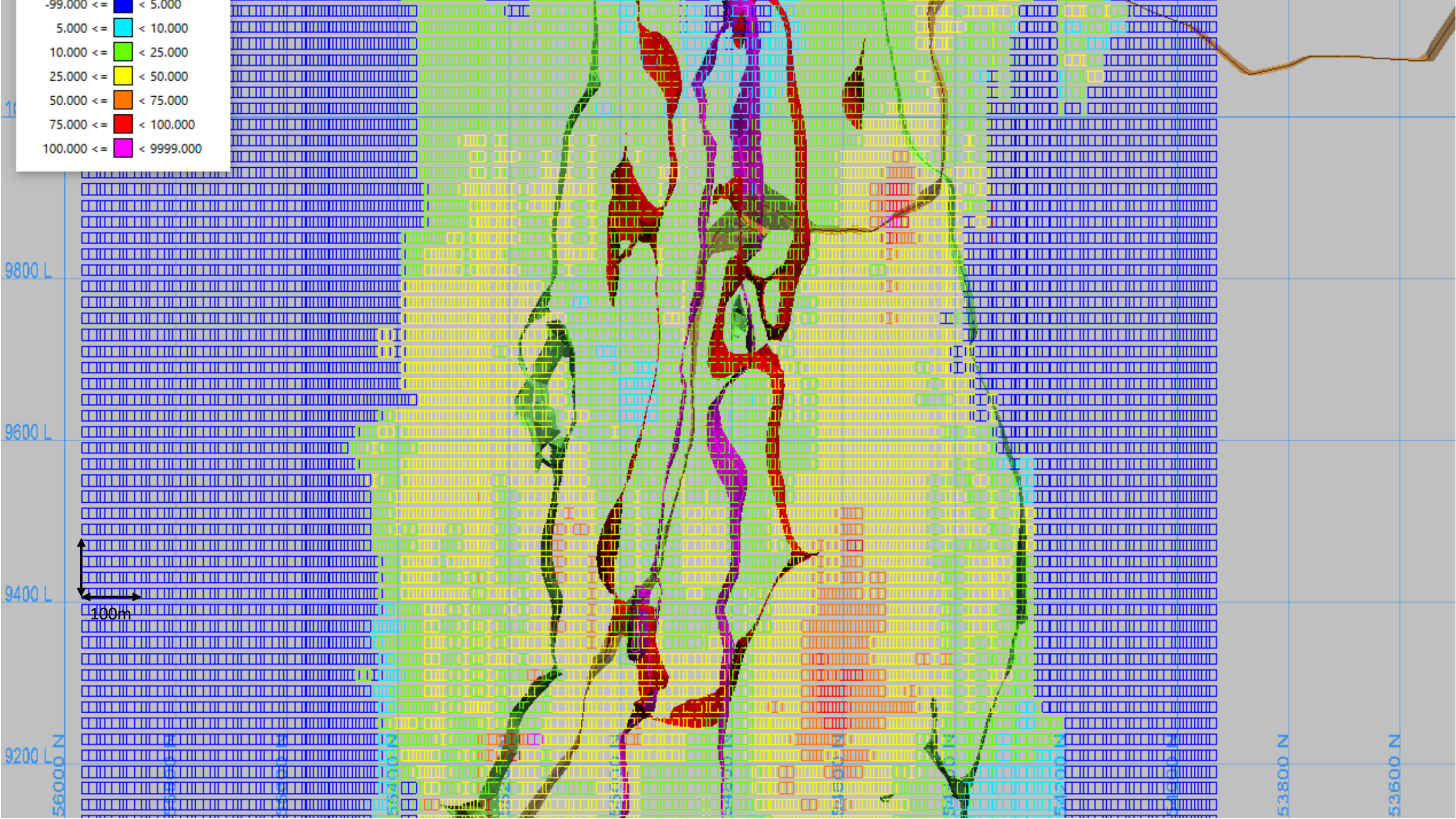
THE GOOD

E26



- Brittle volcanic, and monzonite (BQM) hosts
- Multiple mineralising phases
- >600m of mineralizing QMP emanating through and above BQM stock
- Successive overpressure events on carapaces
- Sulphate-stable system – anhydrite-quartz veins
- Higher vein density associated with higher Cu and Au grades
- Vein-hosted, fracture controlled, and disseminated mineralization
- Au focused in and around QMPs - Mo outboard of best Cu
- Evidence of early fluid leakage – high level phyllic argillic alteration and elevated As, Mo, Sn, W in shallow portions of system
- Average core grades 1.2-1.8% Cu, 0.2-0.6ppm Au

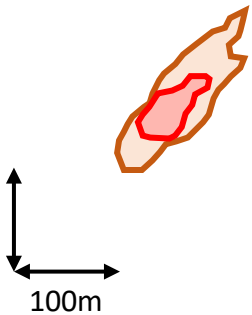




THE BAD

Not As Good

Nerrad



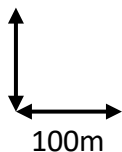
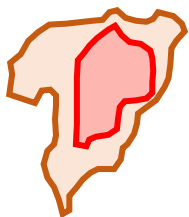
- Massive monzonite (BQM) and early fractionated equivalent (QMZ) hosts
- Single recognized mineralizing phase - entirely enclosed by stock
- Diffuse gradational contacts, common anisotropic textures - Little to no evidence of overpressure
- Sulphate-stable system – Anhydrite & quartz as phenocrysts
- Higher Cu and Au associated with less fractured core zones in mineralizing intrusion
- Disseminated >> fracture controlled>>vein-hosted, mineralisation
- Mo bearing veinlets crosscut strongest mineralization – minor evidence of fluid leakage – generally late quartz-sericite-pyrite faults
- Average core grades 0.6-0.8% Cu, 0.4-2.0ppm Au



AND THE **UGLY**

Still OK

GRP314



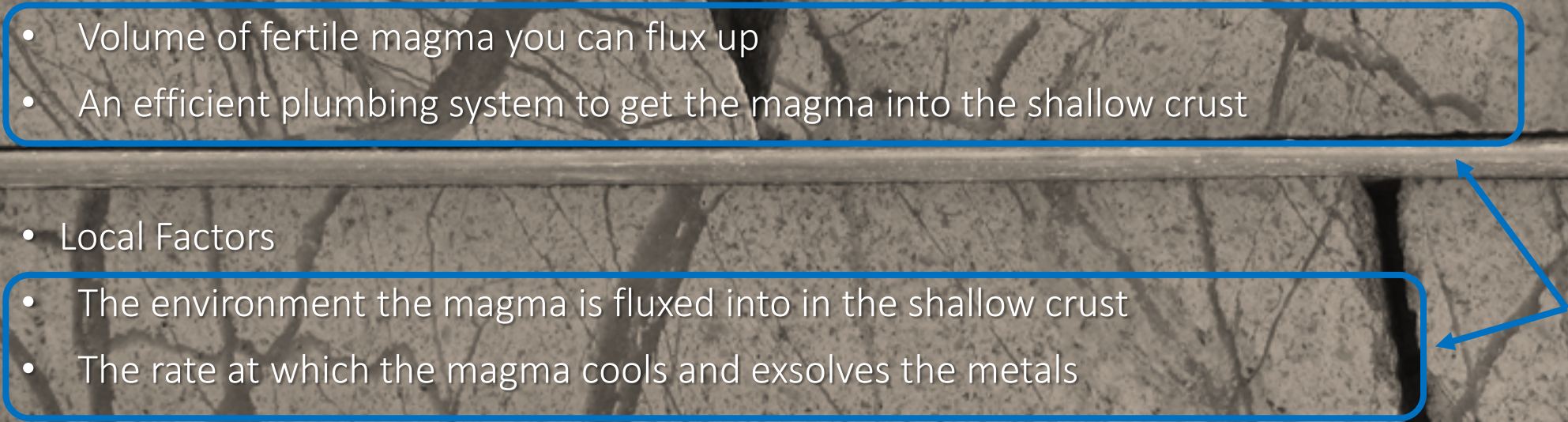
- Hybrid system
- Massive monzonite (BQM) and early fractionated equivalent (QMZ) hosts
- Multiple mineralizing phases - within stock
- Some sharp, some diffuse gradational contacts, local anisotropic textures – some evidence of overpressure
- Sulphate-stable system – anhydrite & quartz as phenocrysts and veins
- Higher Cu and Au associated with stronger veining, lower tenor high-grade zones in less fractured late syn-mineral intrusions
- Fracture controlled >> vein-hosted = disseminated mineralisation
- Mo bearing veinlets over the top of the system, less at depth – common evidence of fluid leakage as late quartz-sericite-pyrite faults
- Average core grades 0.5-0.8% Cu, 0.1-0.3ppm Au

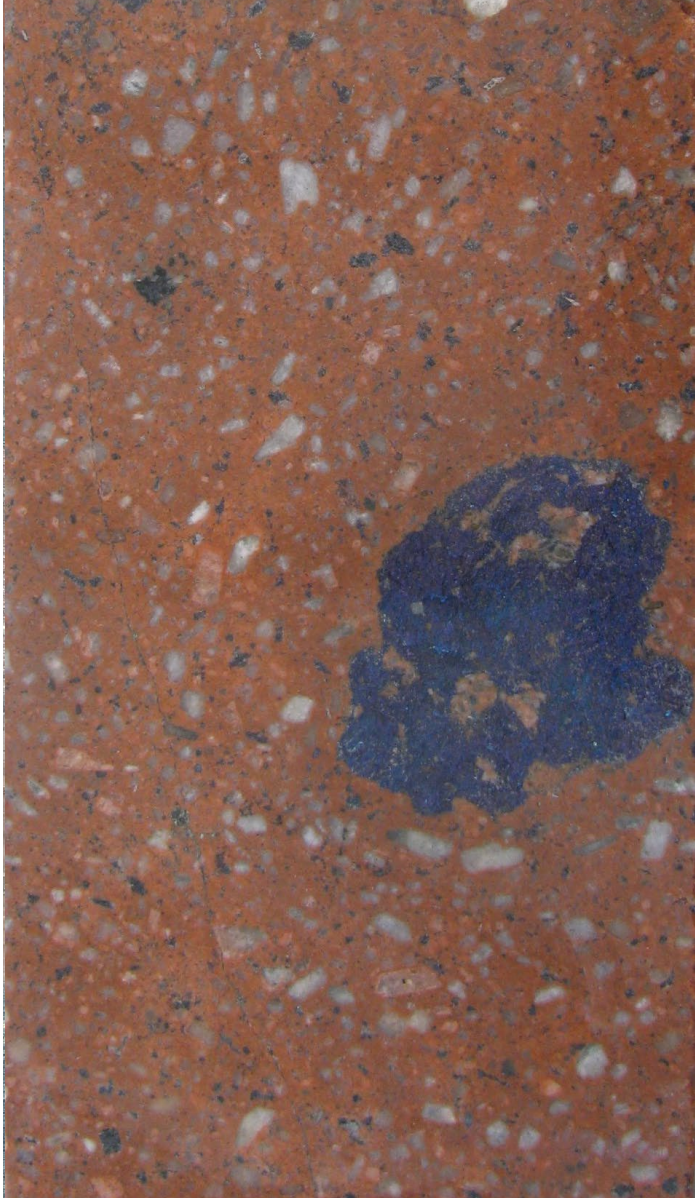


Why the Differences Between the Good, the Bad, and the Ugly?

- Fundamental Controls
 - Fertility of the Magma
 - Volume of fertile magma you can flux up
 - An efficient plumbing system to get the magma into the shallow crust
-
- Local Factors
 - The environment the magma is fluxed into in the shallow crust
 - The rate at which the magma cools and exsolves the metals

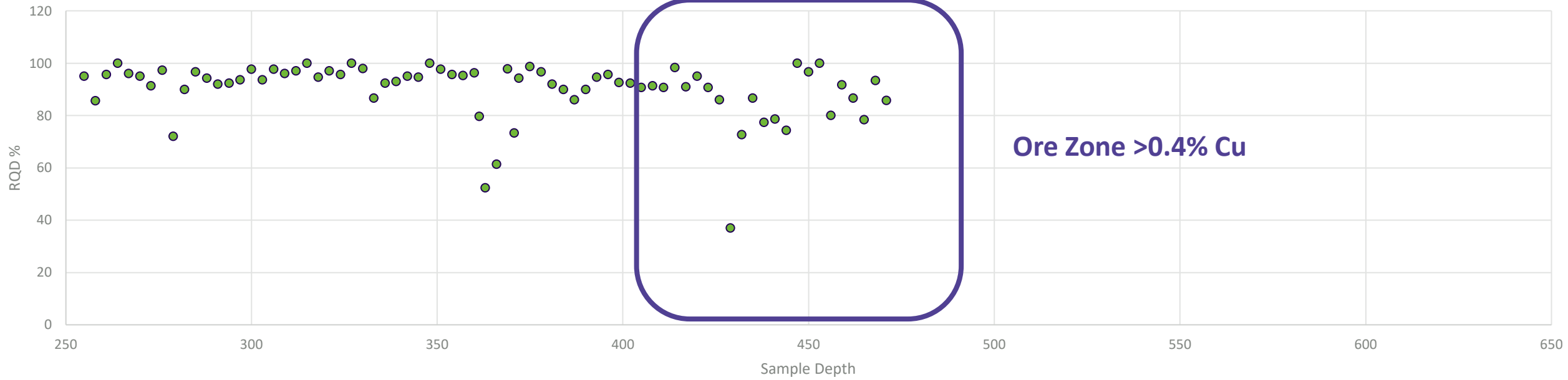
Elements of
Failed
Overpressure
in the system



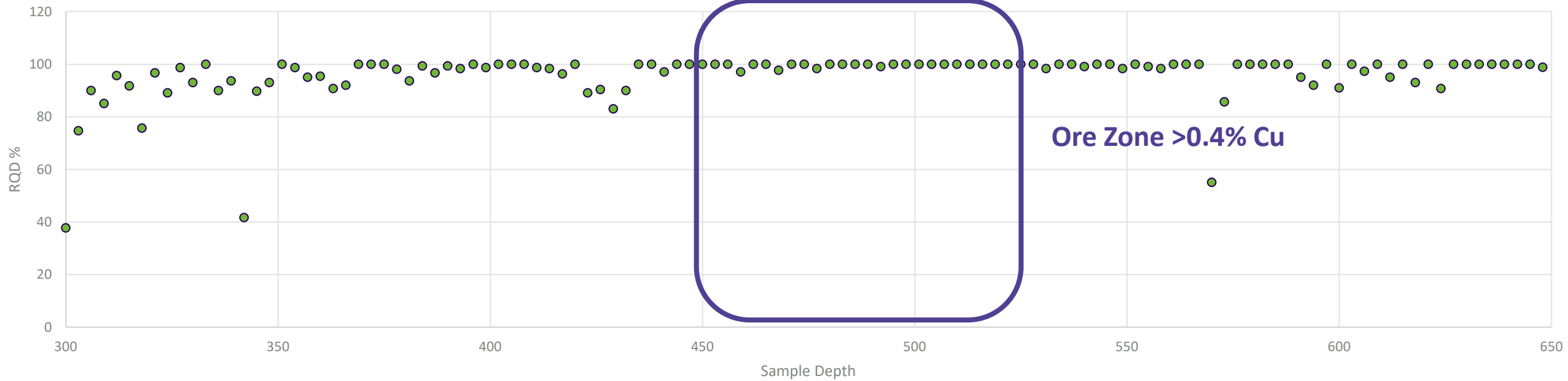


- Confining pressure too great for volume of magma fluxed = stalled system
- *Recognised through texture, contacts, and vein density*
- Absence of clear discordant contacts –small thermal gradient between host / porphyry
- ‘Mushy’ contact breccias, with grain-boundary controlled margins on wallrock clasts
- Abraded feldspar phenocrysts in a f.g. to m.g. equigranular groundmass
- Abundance of quartz or sulphate ‘phenocrysts’
- Lower vein density – low # natural rockmass defects (RQD)
- Dominance of disseminated sulphide –at grain boundaries in groundmass
- Anisotropic textures – mairolitic cavities, sulphide clots
- Strongly telescoped alteration assemblages

E26D559



GD882



Failed Overpressure Due to Depth?

- If Failed Overpressure was solely a function of increased confining pressure at depth...
- Expect to see similar effects at depth in the known economic systems (E26, E48)
- Drilling >1400m vertical in E26, and >1100m vertical in E48- within BQM stock
- Some features present – e.g. ‘mushy’ contacts
- Still strong evidence of overpressure
- Implies differences between the systems due to volume of fertile magma fluxed, and time difference between host / mineralizing porphyry



Implications of Failed Overpressure

- Importance of recognizing local controls to mineralisation
- Variability within Districts/Camps
- Early recognition may give clues to scale of mineralized system and grade
- Also changes the exploration process – alteration and grade may indicate an ‘edge of core’ position, when actually in centre already
- Prospectivity modelling based on relative position to vs. within regional stock
- Different exploration methods – not necessarily looking for ‘damaged rock’



Conclusions

- Significant economic porphyry systems can occur within larger comagmatic intrusions, where a permissive setting is present (e.g. E26 deposit)
- 'Hot' hosts like the BQM and QMZ at Northparkes can lead to gradual broad fluid loss at time of formation, leading to small or lower grade systems, depending on volume of magma batched up into shallow crust
- Hard boundaries on larger stocks, where intruded by mineralising porphyries can inhibit magma ascent – homogenising magma with cooling host rock
- These conditions prevent overpressure from occurring in low-volume mineralising magma batches
- Slow-cooling also inhibits development of typical porphyry features such as true porphyritic texture and stockwork vein arrays formed during overpressure
- Care is required to assess potential at each exploration stage



Acknowledgements

- Past discoverers/workers on the stock hosted deposits: Dedy Hendrawan, Adam Schwarz who set the ball rolling...
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- Thanks to the organising committee for the opportunity, and CMOC – Northparkes Mines for permission to present